

Simulation of Coupled Processes of Flow, Transport and Storage of CO₂ in Saline Aquifers

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Outline

- **Research project and team members**
- **Scope of work**
- **Physical and chemical processes of CO₂ geosequestration**
- **Convective mixing**
- **Rock deformation**
- **Parallel simulation**
- **Future work**

Research Project and Team Members

DOE (NETL):

Area of Interest 2: Simulation

Duration: four year (2009-2013)

DOE Project Manager: Bruce Brown

Colorado School of Mines (CSM):

Yu-Shu Wu, Hossein Kazemi, Xiaolong Yin,
Jeffery Chen, Phil Winterfeld, and Ronglei Zhang

Lawrence Berkeley National Laboratory (LBNL):

Karsten Pruess (LBNL)

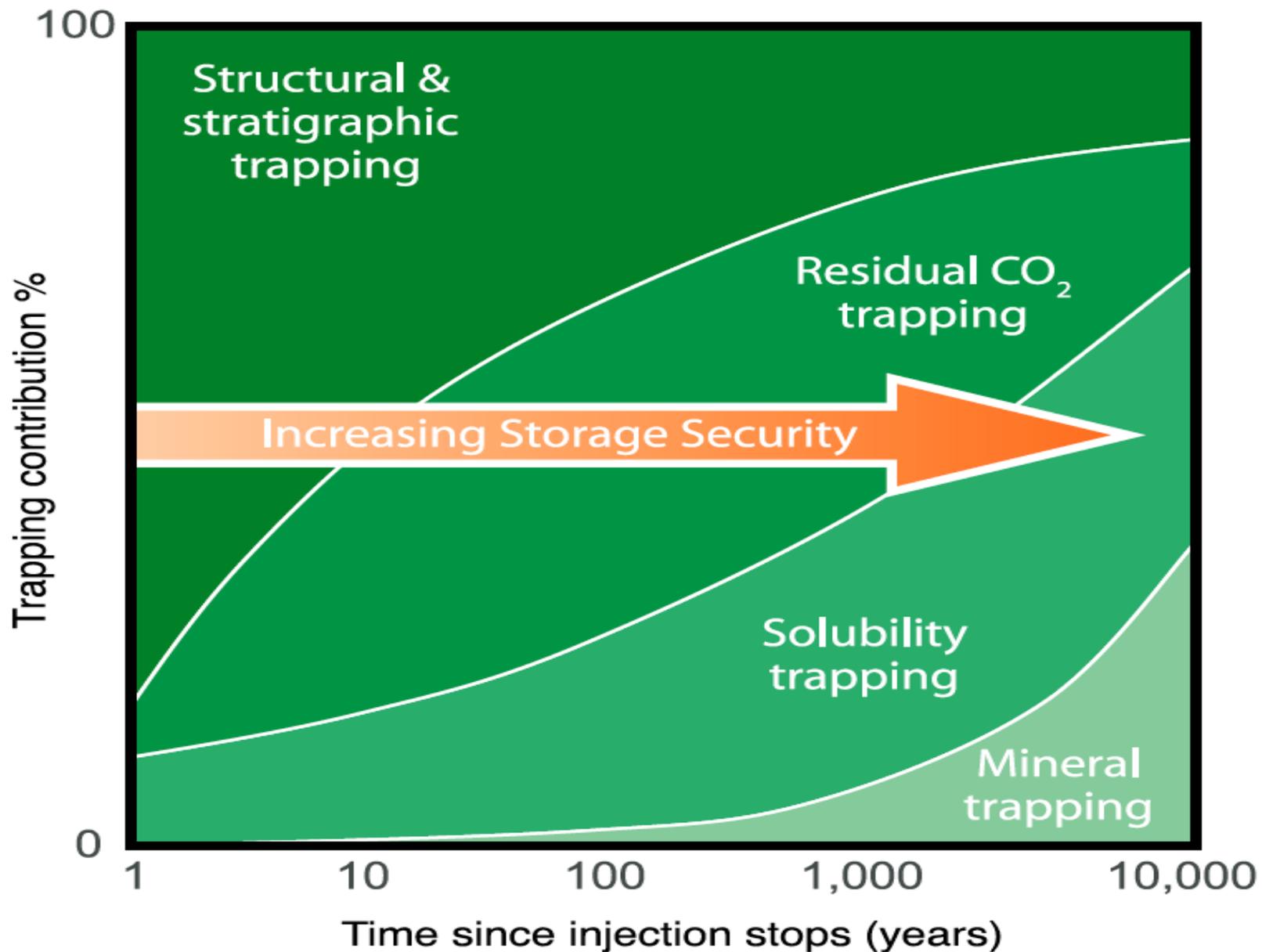
Computer Modeling Group (CMG):

Industry sponsor

SCOPE OF WORK

- 1. Development of a three-phase, non-isothermal flow model to handle CO₂-brine multiphase flow process (Pruess)**
- 2. Coupling of pressure and rock mechanical deformation (Wu and Winterfeld)**
- 3. Incorporation of representative geochemical reactions (Yin)**
- 4. Characterization and modeling of the instability (Yin)**
- 5. Implementation of parallel computing technologies (Chen)**
- 6. Incorporation of general fracture conceptual models (Kazemi and Wu)**
- 7. Model verification and application (All)**

Security of CO₂ Storage vs. Time (IPCC Report 2005)



CO₂ Trapping Scenarios

Structural Trapping (Trapping CO₂ as a mobile “gas”, a supercritical fluid, in reservoirs)

Residual Gas Trapping (Trapping CO₂ as an immobile “gas”)

Solubility Trapping (Trapping of CO₂ as a soluble component in brine)

Mineral Trapping (Conversion of CO₂ into carbonate minerals: Calcite, Dolomite, Siderite, etc.)

CO2 Storage Capacity in North America

<i>Formation Type</i>	<i>10⁹ Tons</i>	<i>%</i>
Saline Aquifers	3,297 - 12,618	91.8 - 97.5
Unmineable Coal Seams	157 - 178	4.4 - 1.4
Mature Oil & Gas Reservoirs	138	3.8 - 1.1
Total Capacity	3,592 - 12,934	100.0

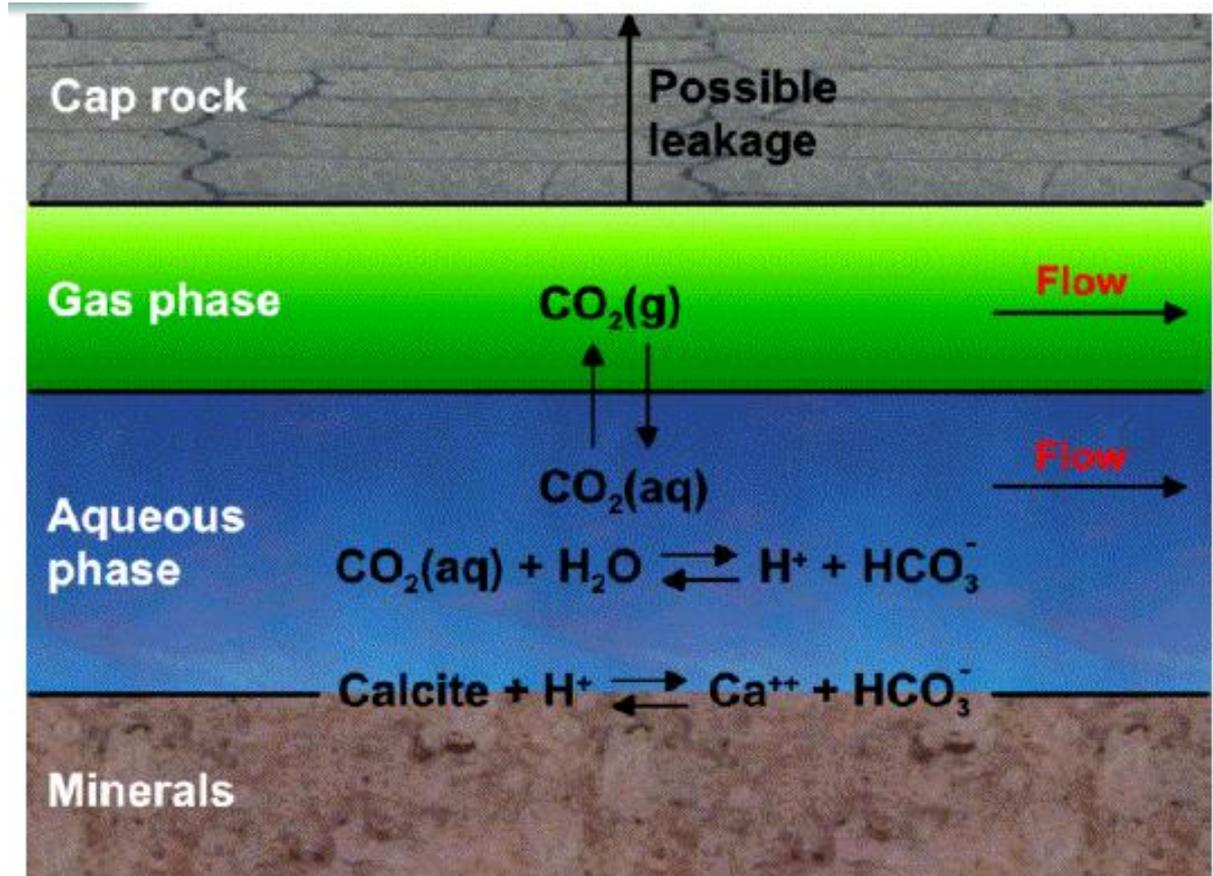
Source: DOE & NETL, "Carbon Sequestration Atlas of the United States and Canada", 2008

Physical / Chemical Processes

Structural Integrity –
Elevated pressure and
effects on existing
fractures / faults

Dissolution – Mass
transfer between CO₂
and formation brine

Geochemistry –
Reactions induced by
dissolved CO₂ with
aquifer rock
(mineralization)

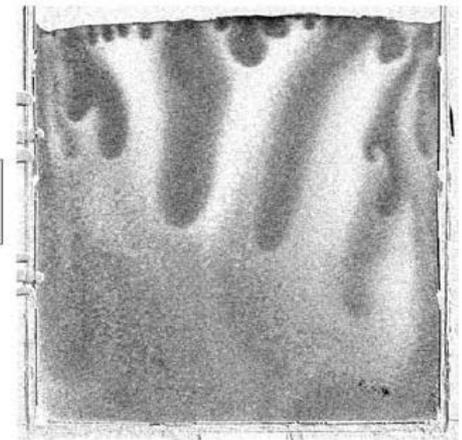
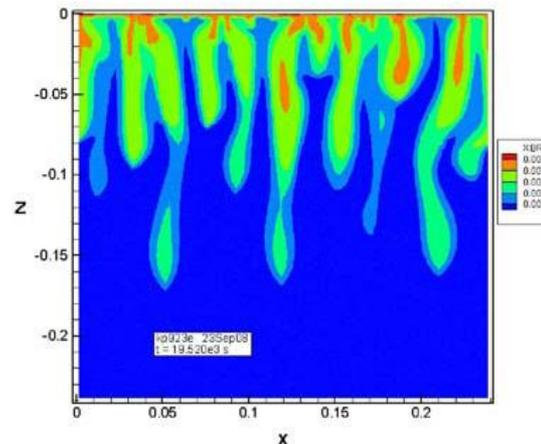
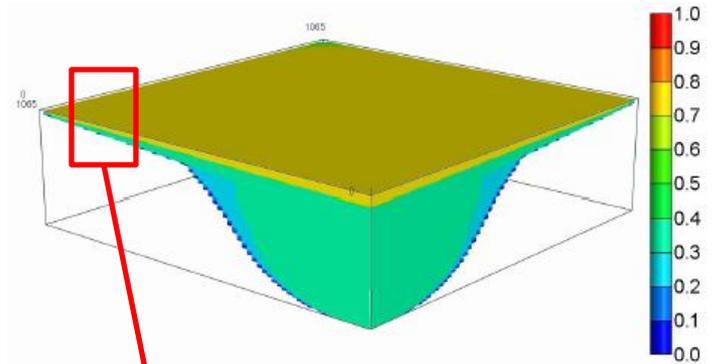


CMG)

Convective Mixing

- Convective mixing is **an instability**
 1. CO₂ plume will migrate to above the brine due to buoyancy
 2. Dissolved CO₂ increases the density of brine below
 3. The heavier brine sinks into the lighter brine below, generating a convection
 4. Convection enhances mixing and solubility trapping

(Adapted from data provided by CMG)

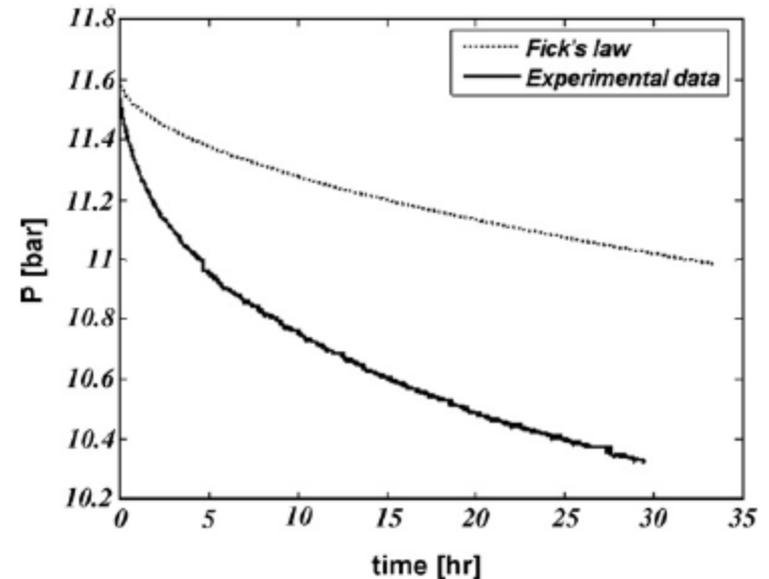


(Ennis-King & Paterson, 2005)

Kneafsey & Pruess, 2010 (Simulation & Experiment)

Significance of Convective Mixing

- Convective mixing leads to **significant increase** in dissolution rate
 - Lower risk of leakage
 - Potentially higher rate of mineralization
- The effect of convective mixing must be properly incorporated into CO₂ sequestration simulators



Farajzadeh et al. 2007 – Pressure reduction measured in experiment vs. calculation based on Fick's diffusion

Integral Properties and Modeling

- Fick's law of diffusion

$$\mathbf{q}_c = -\underline{\underline{\mathbf{D}}}\nabla c$$

- Enhanced diffusion due to convective mixing

$$\mathbf{q}_c^* = -\underline{\underline{\mathbf{D}}}\nabla c - D_{cm} \frac{\partial c}{\partial z} \mathbf{e}_z$$



We will use 2D and 3D simulations to characterize D_{cm} as a function of permeability, anisotropy, heterogeneity, and concentration gradient

Mechanical Coupling:

Effect of Pressure and Temperature on Rock Deformation

Poroelasticity and effective stress:

$$\varepsilon_b = \frac{1}{K} P_c - \alpha P_p = \frac{1}{K} \sigma'$$

(ε_b is bulk volumetric strain, P_c is confining pressure,
 σ' is effective stress)

Modification of porosity, permeability and capillary pressure due to rock deformation

Effective Stress Dependence

- Mckee et al. (1988) – porosity depends on bulk rock compressibility (c_m) and effective stress, Carman-Kozeny equation for permeability:

$$d\phi = -c_m \frac{\phi^3}{1-\phi} d\sigma', \quad k \propto \frac{\phi^3}{1-\phi^2}$$

- Ostensen (1986) – permeability correlation for tight gas sands:

$$k^{0.44} = C \ln \left[1 + \left(\frac{\sigma^*}{\sigma'} \right)^2 \right]$$

Simulation Example

- Modified example from Rutqvist et al. (2002)
- Brine formation (200 m thick)
- 2D system, large x-direction length and constant pore and confining pressure gradients
- CO₂ injection at one end, formation bottom, of 100x5 x-z grid for ten years
- Permeabilities and porosity calculated using fractured rock correlations (Rutqvist et al. 2002)
- Pc scaling using J-function

Comparison with Rutqvist et al. Gas Saturations

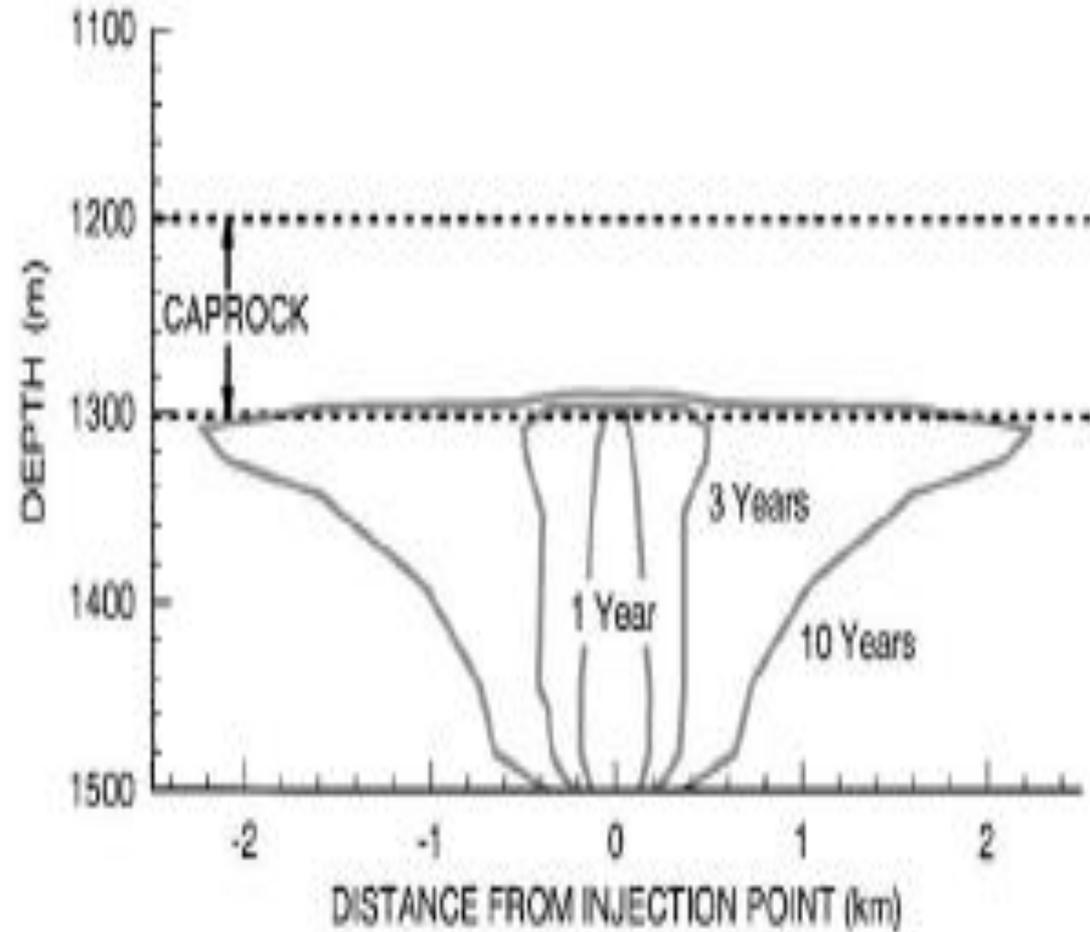
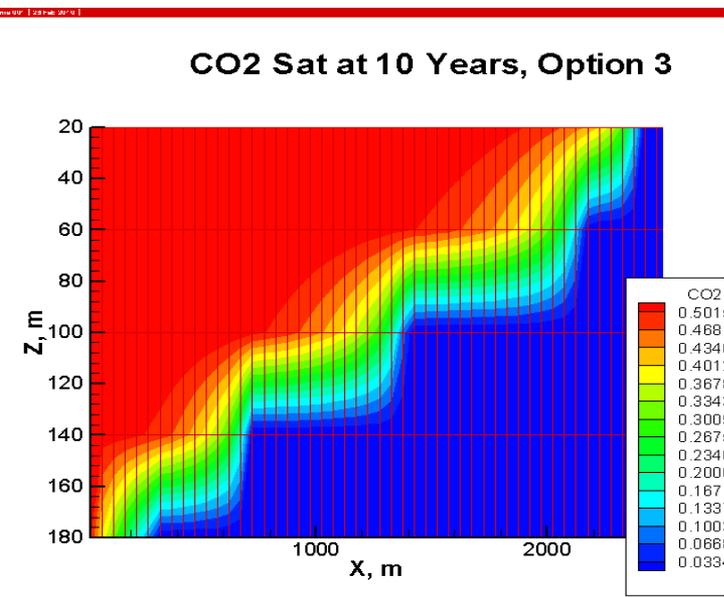
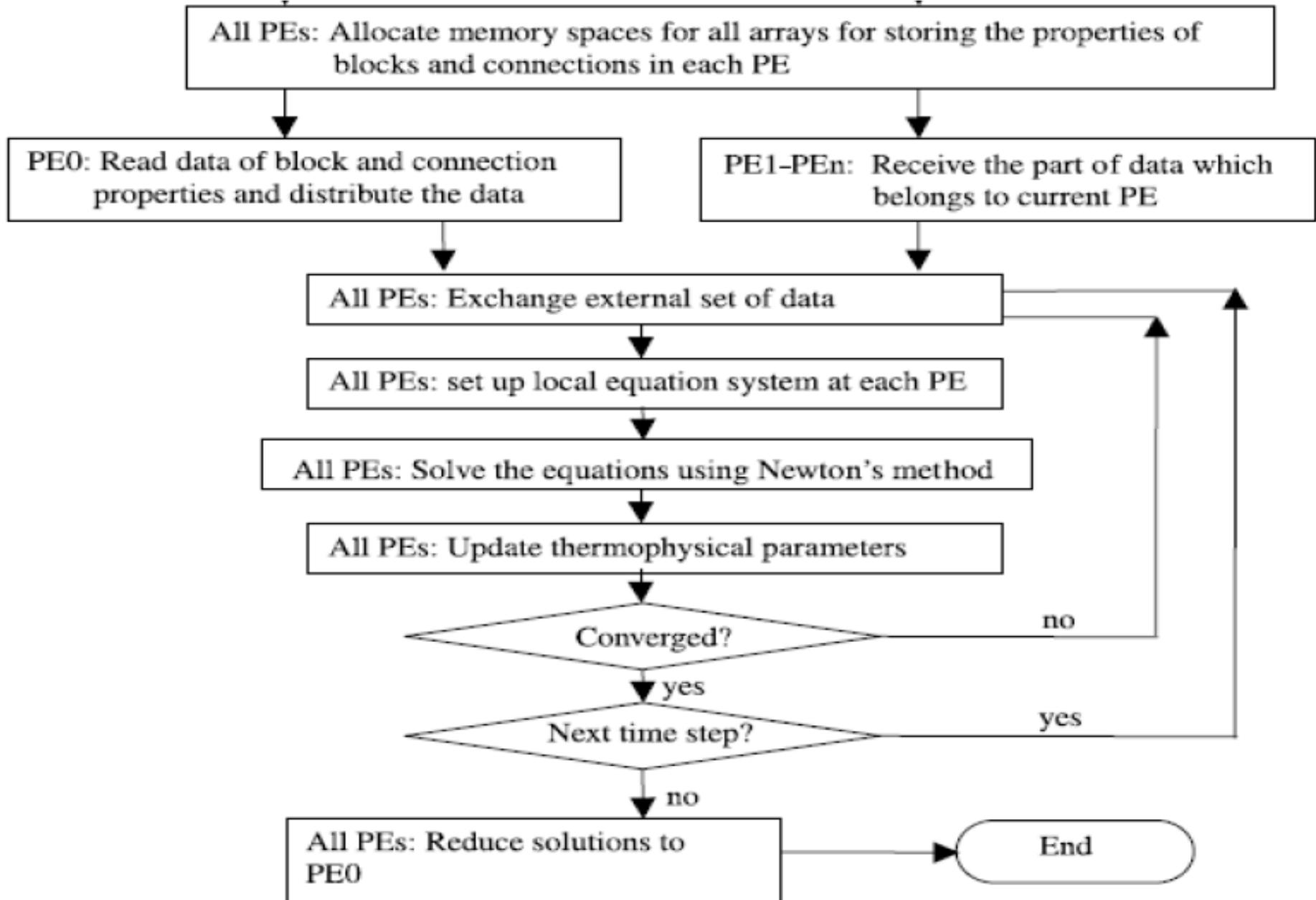


Fig. 6. Spread of the supercritical CO₂ fluid within the brine formation.

Parallel Simulation



Future Work

1. Develop and implement a three-phase flow module
2. Coupling of pressure and temperature effect on rock mechanical deformation
2. Integration of kinetic modeling of mineral dissolution and precipitation and non-aqueous reactions into models
3. Characterization and modeling of the instability
4. Improvement of parallel computing technologies
5. Development and incorporation of fracture conceptual models

Thanks!