



Transport and Dispersion Processes for CO₂ in the Unsaturated Zone and Surface Layer

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Introduction and Motivation

- Injection of CO₂ into deep geologic formations involves risk that CO₂ will migrate away from primary target.
- We are carrying out research on CO₂ transport for:
 - (1) assessment of health, safety, and environmental (HSE) risk;
 - (2) development of sequestration verification approaches (instrumentation requirements and monitoring strategies).
- CO₂ seepage flux and near-surface concentration are risk drivers.

Approach:

Coupled subsurface-surface layer simulation
to estimate CO₂ fluxes and concentrations

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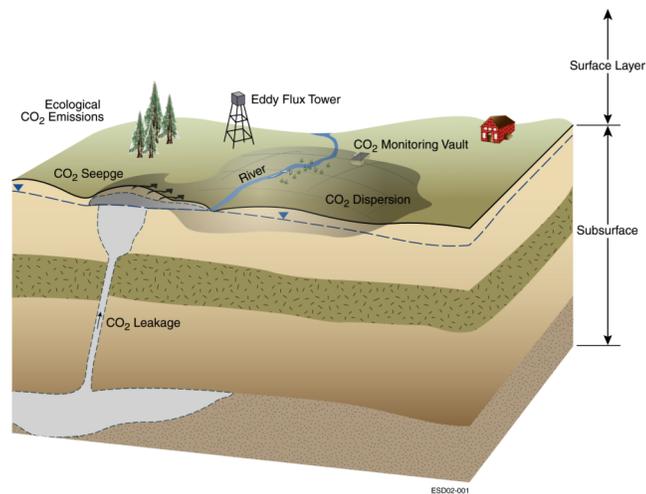
Outline



- Terminology.
- Leakage and seepage processes.
- Seepage case studies.
- Process modeling.
- Physical properties.
- Unsaturated Zone (UZ) simulations.
- Coupled UZ–surface layer simulations.
- Conclusions.

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Leakage and Seepage from Geologic Carbon Sequestration Sites



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Terminology



- Leakage = CO₂ migration away from primary sequestration target.
- Seepage = CO₂ transport out of the ground into the atmosphere or into surface water.
- Leakage/Seepage can be diffusive and/or advective.
- Dispersion = CO₂ dilution by advective and diffusive processes.
 - Hydrodynamic dispersion in subsurface.
 - Atmospheric dispersion above ground.
- Surface layer = bottom 1/10 of the atmospheric boundary layer.

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How Will CO₂ Leak?



- Upward buoyancy driving force for CO₂ is present in saturated rock. Geothermal gradient ensures positive (upward) buoyancy force.
- The length scale for leaking CO₂ increases:
 - Buoyant CO₂ will spread out laterally against permeability barriers.
 - CO₂ decompresses as it migrates upwards.
- Permeability is a scale-dependent property. Higher-k pathways more likely to be found by a larger buoyant and spreading plume.

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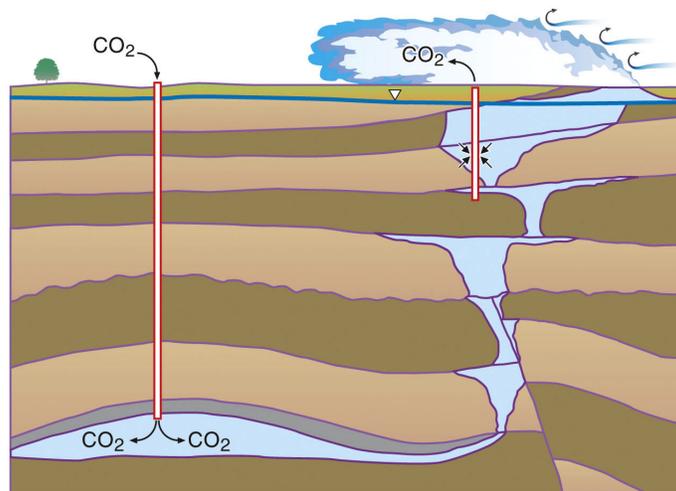
How Will CO₂ Seep?



- CO₂ will convectively seep by upward buoyant flow if no unsaturated zone.
- CO₂ will advectively seep if there is a continuous source from the saturated zone.
- CO₂ will diffusively seep if concentrations exceed soil and ambient air concentrations.

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CO₂ Leakage, Seepage, and Dispersion



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Case Studies: Seeping Gases



- Numerous examples of seeping gases exist.
 - Natural gas vents and blow holes (e.g., Rio Vista, ca. 1900).
 - Trace hydrocarbon gases above oil fields (e.g., Las Animas Arch, KS-CO).
 - Natural gas seeps (e.g., Santa Barbara Channel, CA).
 - Accidental natural gas storage releases, lateral migration, and escape through abandoned wells (e.g., Hutchinson, KS, Leroy Gas Storage, WY).

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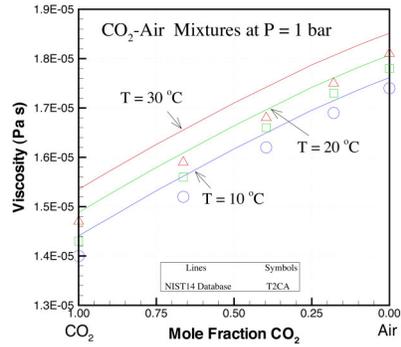
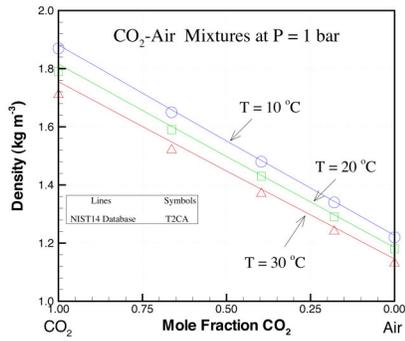
Process Modeling Approach



- T2CA (TOUGH2 for CO₂ and Air)
 - water, brine, CO₂, gas tracer, air, heat.
 - real gas properties for ρ and μ .
 - Henry's Law for solubility.
 - Subsurface and surface layer domains.
 - 3-D integral finite difference method.
- Surface Layer
 - Logarithmic velocity profile (passive flow).
 - Pasquill-Gifford and Smagorinski Model dispersivities.
 - Subsurface and surface layer are coupled.

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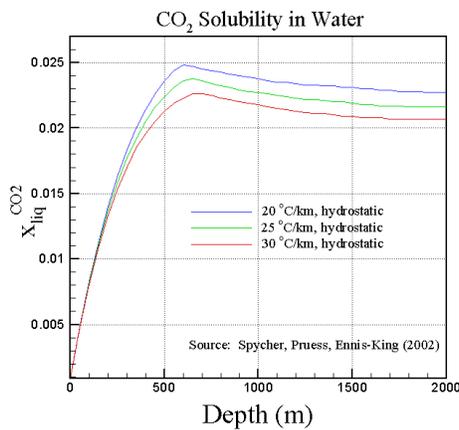
Near-Surface CO₂ Density and Viscosity



T2CA density and viscosity models agree to within 5% of NIST14 Database values.

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Gas Solubility in Water



T = 25 °C, 1 atm

Solubility (mol L⁻¹)

N₂ 0.68 x 10⁻³

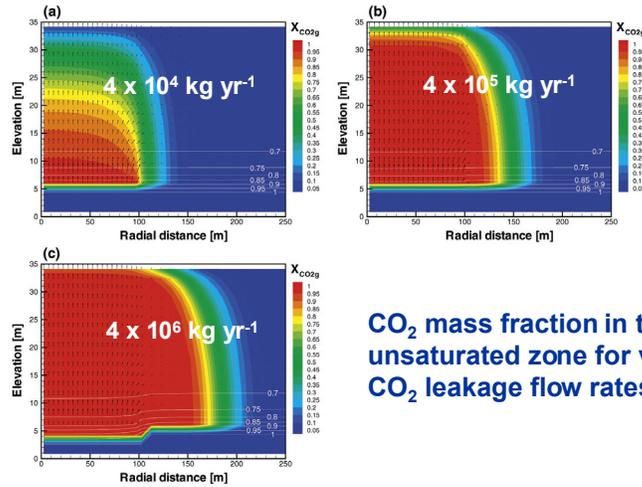
O₂ 1.4 x 10⁻³

CO₂ 32. x 10⁻³

CO₂ is approx. 50x more soluble than air in water.

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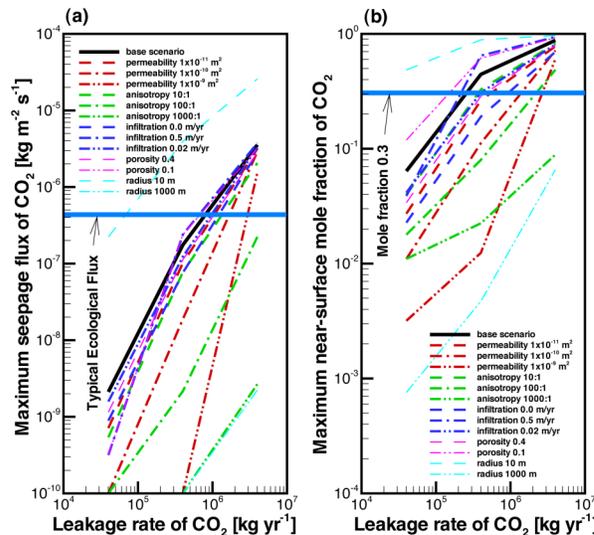
CO₂ Migration Through 30 m of Unsaturated Zone with Infiltration 10 cm yr⁻¹



CO₂ mass fraction in the unsaturated zone for various CO₂ leakage flow rates.

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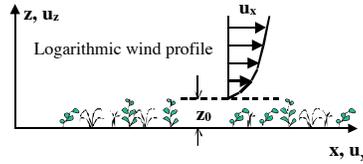
Sensitivity of CO₂ Seepage and Soil Gas Concentration to Various UZ Properties



(Oldenburg and Unger, *Vadose Zone Journal*, 2003)

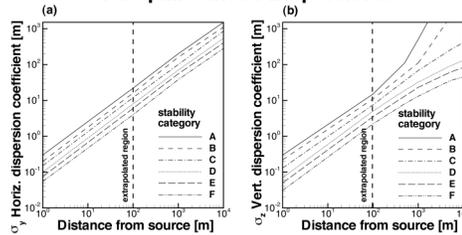
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Surface Layer Modeling



$$u_x(z) = \frac{u_*}{k_*} \ln\left(\frac{z}{z_0}\right)$$

Pasquill-Gifford Dispersivities



$$\sigma_x^2 = 2D_{xx}t$$

$$\sigma_y^2 = 2D_{yy}t$$

$$\sigma_z^2 = 2D_{zz}t$$

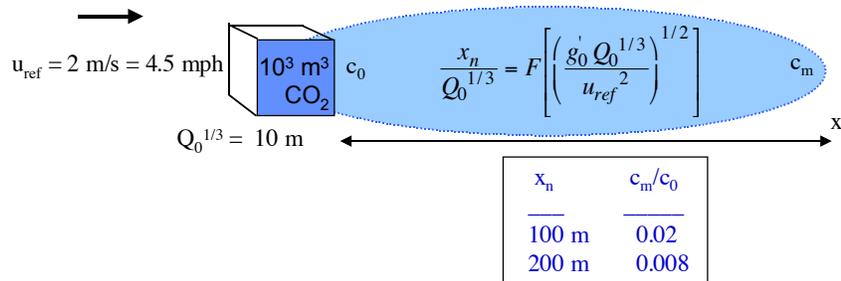
$$D = \frac{l^2}{\sqrt{2}} \left(\frac{\partial u_x}{\partial z} + \frac{\partial u_z}{\partial x} \right) = \frac{l^2}{\sqrt{2}} \frac{u_*}{k_*} \frac{1}{z}$$

Smagorinski Model dispersivities
(*l* is a grid-related length scale).

Empirical Correlations of Atmospheric Dispersion of CO₂



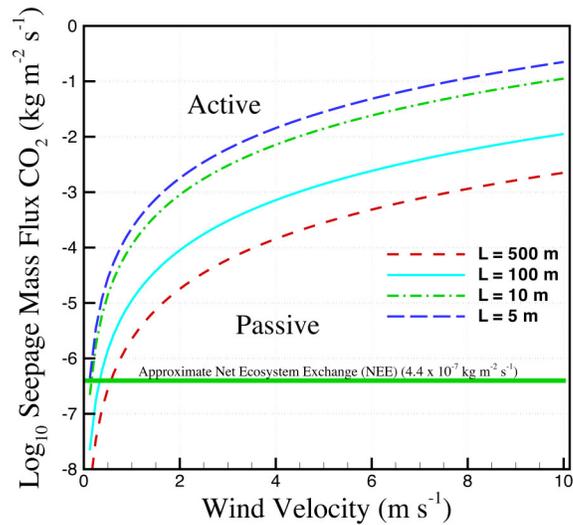
(Britter and McQuaid, 1988)



Empirical results suggest atmospheric dispersion is effective in diluting dense gases over relatively short time and length scales.

However, calm conditions and topographic effects may alter these predictions and require numerical modeling approaches to assess.

Criterion for Active vs. Passive Dispersion for Various Source Length Scales



(correlation of Britter and McQuaid, 1988)

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Properties of Coupled System



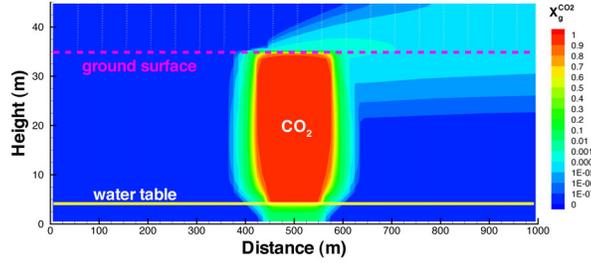
Property	Value
<i>Subsurface</i>	
Permeability ($k_r = k_z$)	$1 \times 10^{-12} \text{ m}^2$
Porosity (ϕ)	0.2
Infiltration rate (i)	$10. \text{ cm yr}^{-1}$
Residual water sat. (S_{lr})	0.1
Residual gas sat. (S_{gr})	0.01
van Genuchten (1980) α	$1 \times 10^{-4} \text{ Pa}^{-1}$
van Genuchten (1980) m	0.2
<i>Surface Layer</i>	
Friction velocity for $u_x = 1 \text{ m s}^{-1}$	0.0869 m s^{-1}
Friction velocity for $u_x = 5 \text{ m s}^{-1}$	0.434 m s^{-1}
Reference height (z_0)	0.10 m
Reference velocity at $z = 10 \text{ m}$	1 or 5 m s^{-1}

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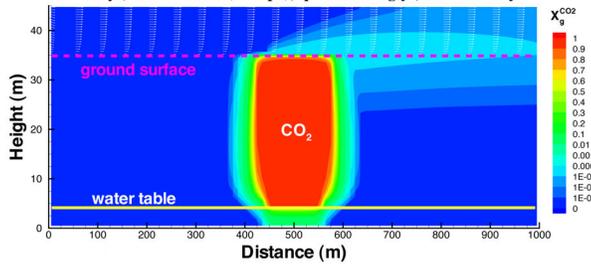
Preliminary T2CA Results



t = 0.5 yr, wind = 1 m/s (2.2 mph), q = 4 x 10⁶ kg/yr, inf = 10 cm/yr



t = 0.5 yr, wind = 5 m/s (11 mph), q = 4 x 10⁶ kg/yr, inf = 10 cm/yr

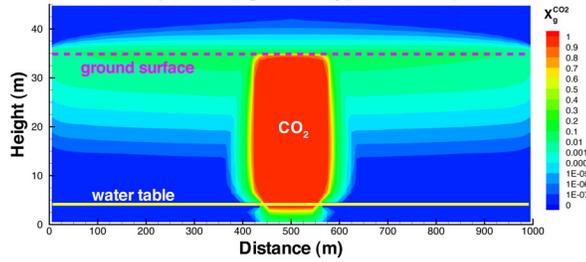


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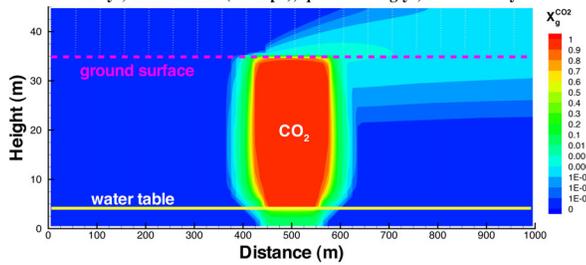
Preliminary T2CA Results



t = 0.5 yr, no wind, q = 4 x 10⁶ kg/yr, inf = 10 cm/yr



t = 0.5 yr, wind = 1 m/s (2.2 mph), q = 4 x 10⁶ kg/yr, inf = 10 cm/yr



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Conclusions



- Leakage and seepage of gases are common.
- Leakage rate and source area strongly control seepage.
- Although diffuse seepage flux may be small, CO₂ concentrations in shallow soil can be large.
- Unsaturated zone can attenuate small leakage fluxes but should not be counted on to attenuate large CO₂ leakage fluxes.
- Rainwater infiltration returns CO₂ to the subsurface.
- Atmospheric dispersion is effective at diluting gases.
- CO₂ concentrations may be higher in low-lying and stagnant (low-wind) areas.

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